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# The effect of mesoscale eddies on air-sea interactions — a view from the ocean

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# Ocean mesoscale eddies

- They are the “**weather**” of the ocean, with horizontal scales of  $\sim 100\text{km}$  and timescales of  $\sim 1\text{month}$ ;
- They typically exhibit different properties to their surroundings, allowing them to **transport properties such as heat, salt and carbon** around the ocean;
- The largest scale eddies emerge from instabilities of strongly horizontally sheared motions, particularly in boundary currents. These eddies often take the form of **well defined rings extending to great depth**. At slightly smaller scales, eddies are generated by baroclinic instability.

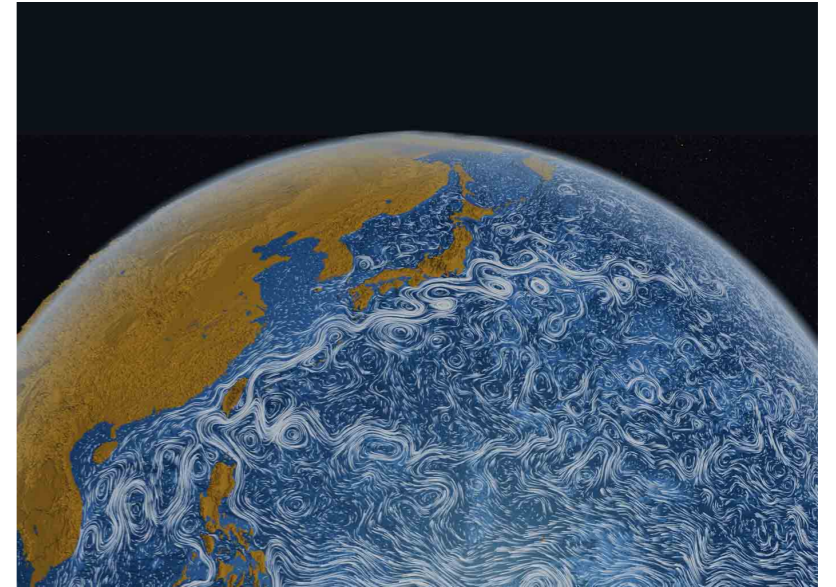


Fig1: NASA visualization of ocean eddies

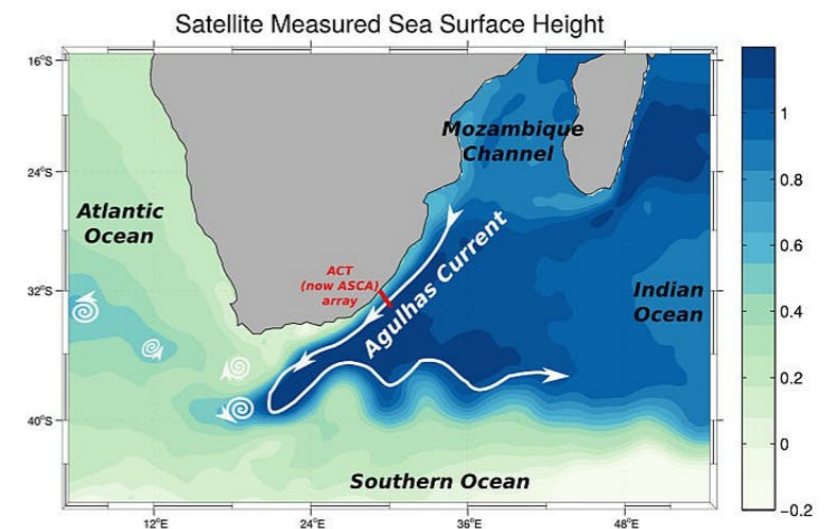


Fig2: Generation of Agulhas rings

# Air-sea interaction over ocean eddies

- At large scales, intensified winds cool the ocean surface through evaporation, as well as increasing the entrainment of upper-thermocline waters into the mixed layer.
- So the **correlation between wind speed and SST** is often **negative**, which means the atmosphere is driving an ocean response.
- However, recent observations show a **positive** correlation over ocean mesoscale features, which is suggestive of the ocean forcing the atmosphere, mainly through heat fluxes out of the ocean.

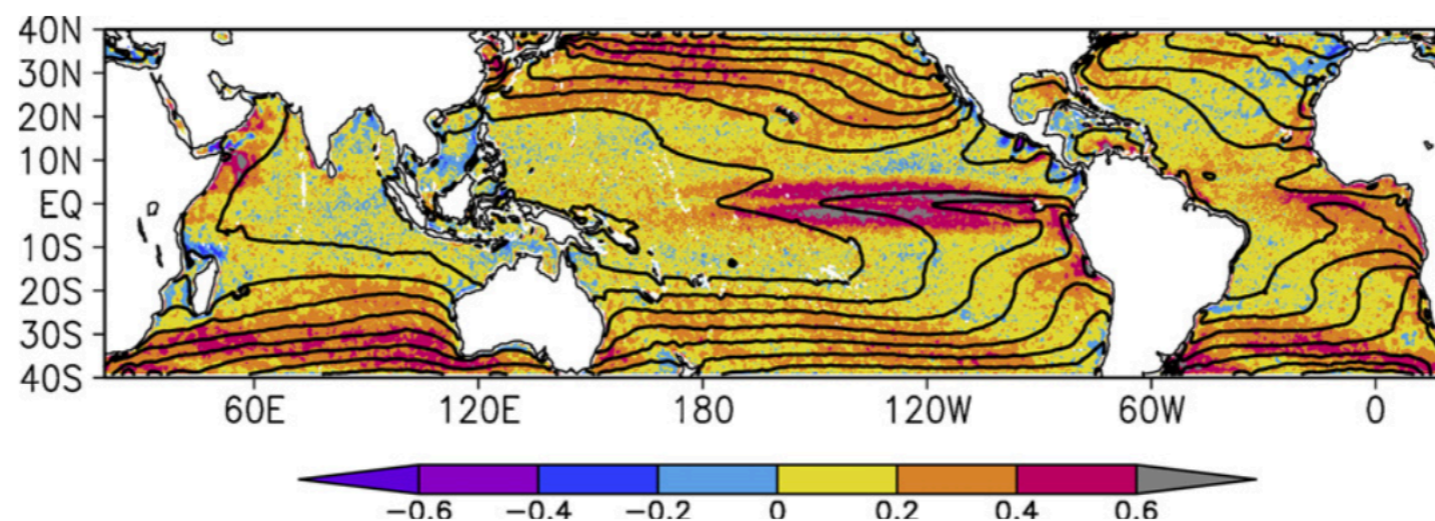
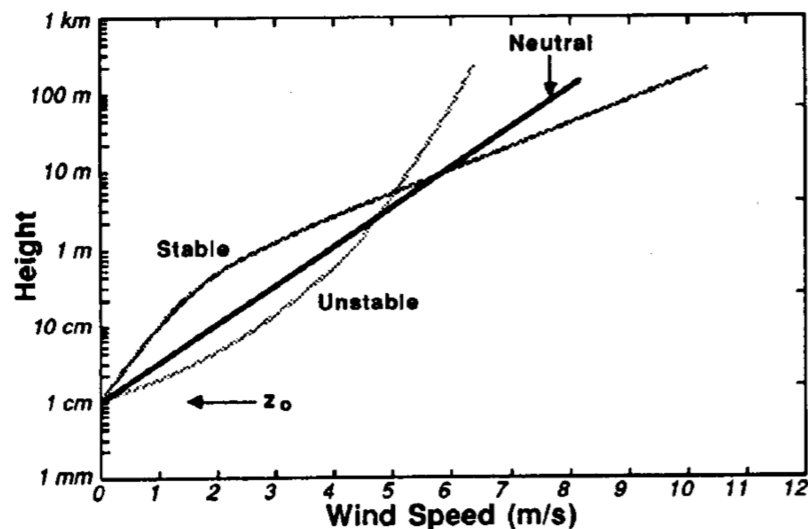


Fig3:A map of the correlation between SST and neutral 10m wind speed at the mesoscale.  
(from R.J. Small et al. 2008)

# Air-sea interaction over ocean eddies (an emphasis on the **atmosphere**)

- There are several reasons why the **atmosphere is affected by ocean eddies**.
- Example 1: as air is blowing across an SST gradient, an air-sea temperature/humidity difference is generated. This leads to changes in **near-surface stability**.



**Vertical gradients of wind speed:  
Stable: increase  
Unstable: decrease**

Fig4: Typical wind speed profiles vs. static stability in the surface layer.  
(from R.B. Stull 1988)

- Example 2: The atmospheric pressure also changes, leading to a spatial pressure gradient which can drive secondary circulations.



# Air-sea interaction over ocean eddies (an emphasis on the **ocean**)

- The viewpoint from ocean response:
  - 1) Surface currents of ocean eddies will impact the relative motion of the air and ocean, acting to change the surface stress, thus affecting the atmosphere as well as **feeding back onto the ocean**.
  - 2) Ocean Ekman layer dynamics is also modified by eddying currents, even with an invariable wind stress.

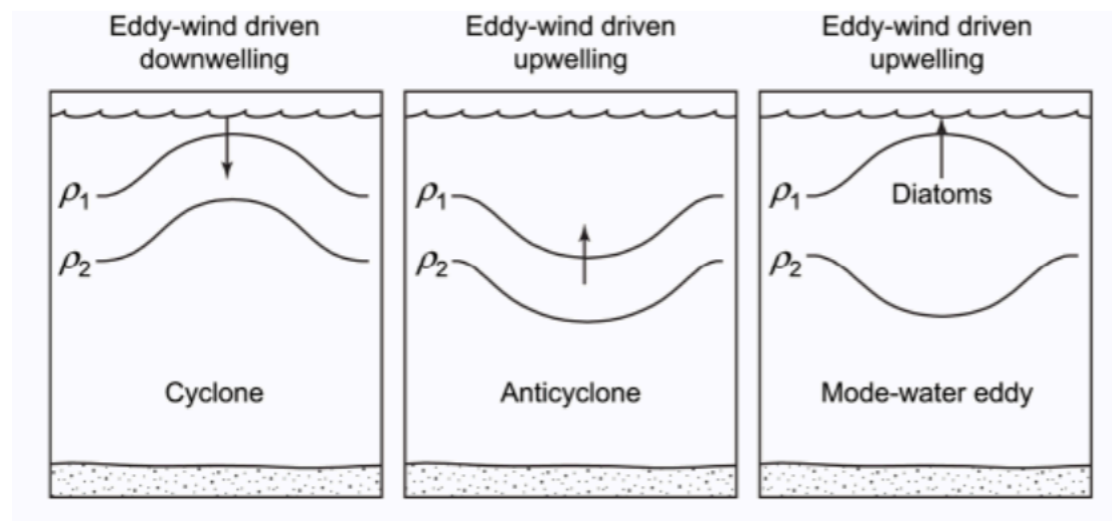


Fig5: Isopycnal displacements associated with three types of eddies. Two density surfaces are depicted seasonal and main thermoclines respectively.  
(from D.J. McGillicuddy et al 2007)

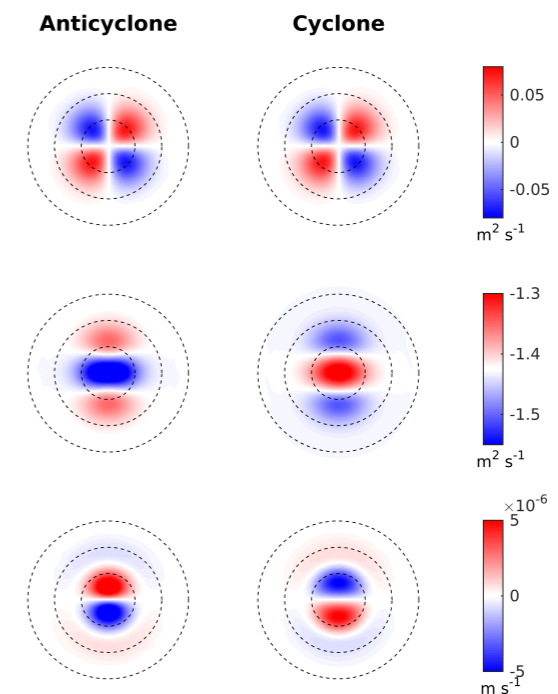


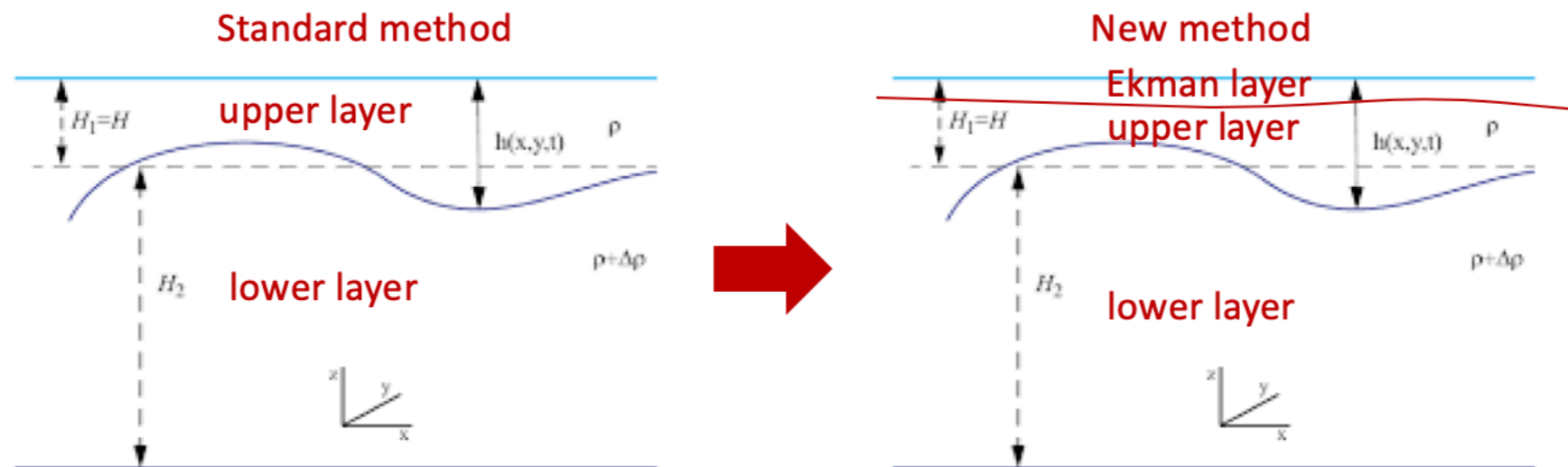
Fig6: Eddy-dependent zonal transport, meridional transport and Ekman pumping velocities.

# Flow-dependent Ekman dynamics

- The Ekman equation in our model

$$\frac{\partial \mathbf{u}_{Ek}}{\partial t} + (\mathbf{u}_{Ek} \cdot \nabla) \mathbf{u}_0 + (\mathbf{u}_0 \cdot \nabla) \mathbf{u}_{Ek} + (\mathbf{u}_{Ek} \cdot \nabla) \mathbf{u}_{Ek} + f \hat{z} \times \mathbf{u}_{Ek} = \frac{\partial \boldsymbol{\tau}}{\partial z} - A_h \nabla^4 \mathbf{u}_{Ek}$$

- Comparison between two regimes:
  - Standard: wind stress is applied as a body force in the upper-layer momentum equation;
  - New: use an explicit Ekman layer to force the upper-layer mass equation.



# Flow-dependent nonlinear Ekman dynamics

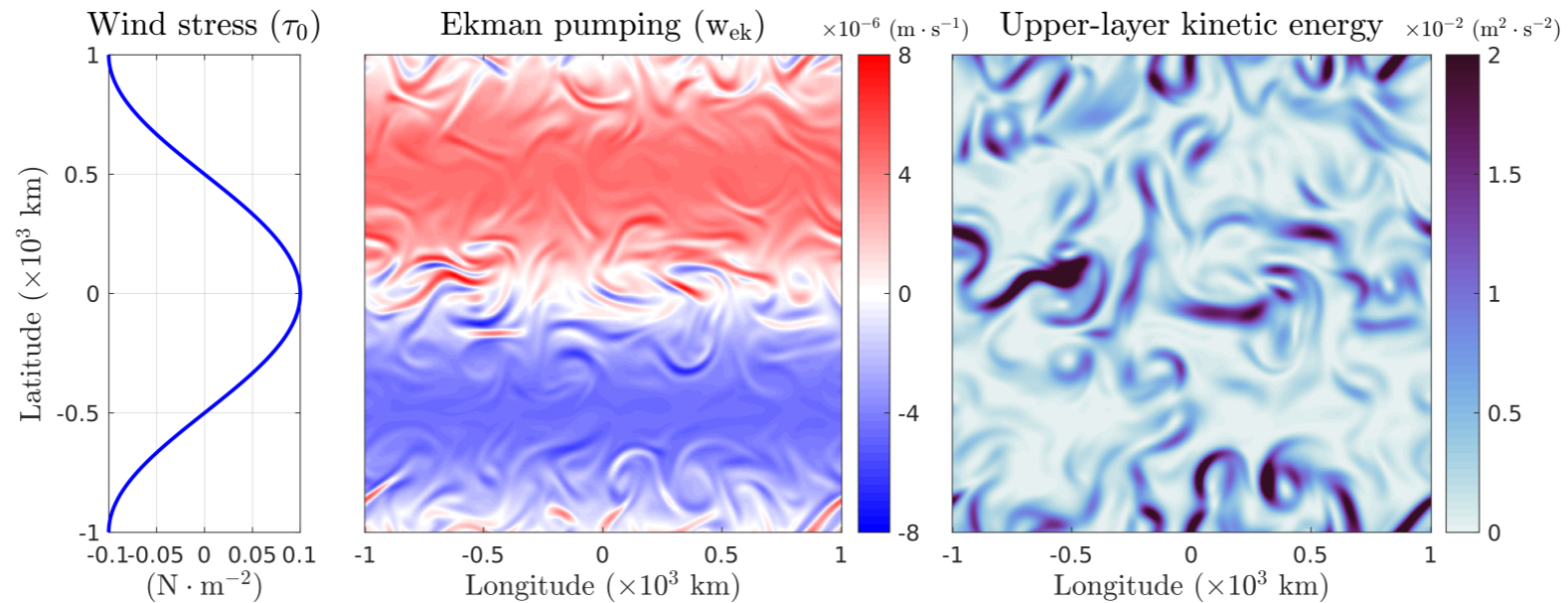


Fig7: The two-layer shallow water model response to a baroclinically unstable jet (with the new method).

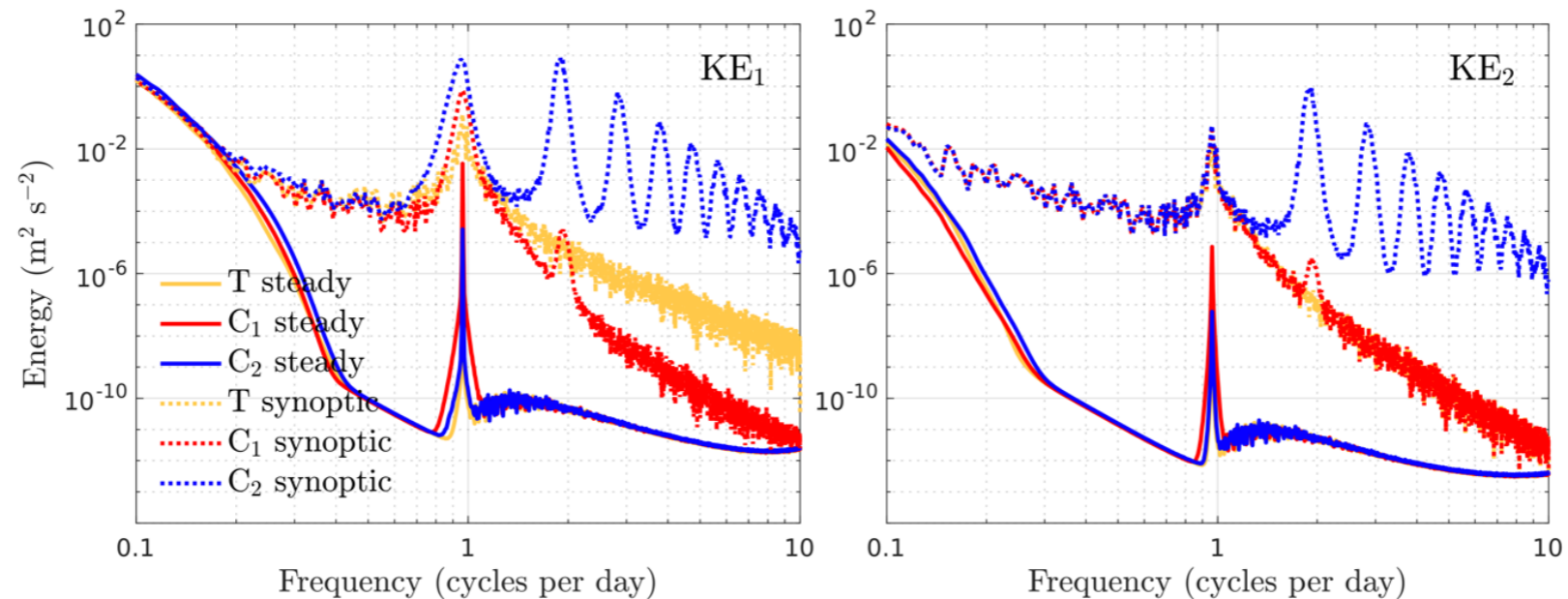


Fig8: The frequency spectra of upper and lower layer kinetic energy (T is the standard two-layer model; C1 and C2 are the new coupled models).

# EUREC<sup>4</sup>A and ocean eddies

- Very few observations are available to quantify the role of ocean eddies in the transport of water properties and in air-sea interactions, especially in the tropics.
- In particular, intense warm eddies converge in the western tropical Atlantic, offshore of Barbados. They carry freshwater from the Amazon river, which results in intensification of storms.
- Moreover, a synoptical study from different research vessels measuring different mesoscale eddies across the experimental area will provide new information on water-mass characteristics advected by the regional eddies. Vertical profiling of the water column will be measured as well.

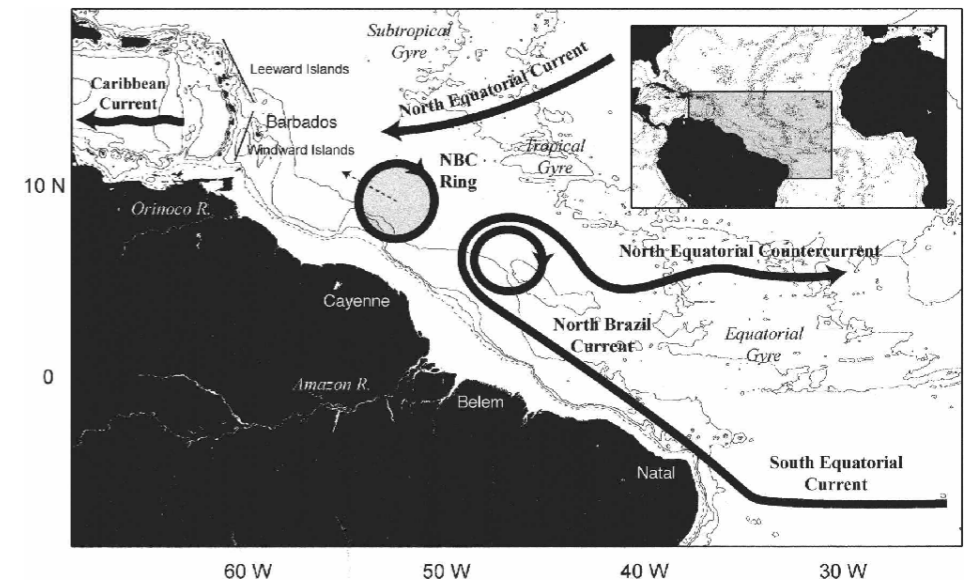


Fig9: Circulation in the western tropical Atlantic Ocean (from D.M Fratantoni et al. 2006).

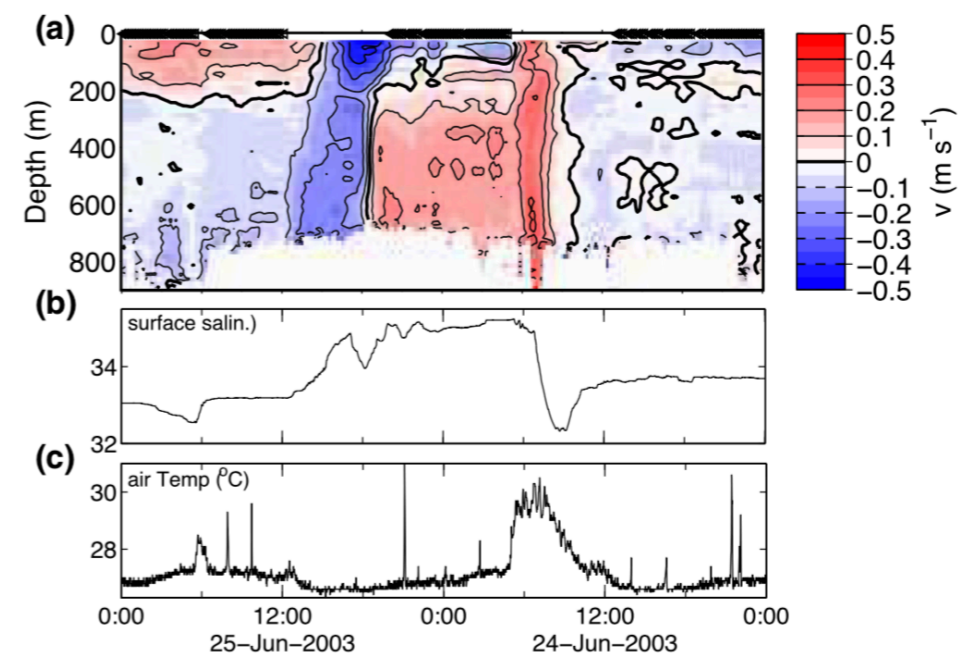


Fig10:RV SONNE ship survey through an anticyclonic eddy west of Barbados.  
(a)Meridional current ; (b)Sea surface salinity; (c) Air temperature.



**Thanks for your attention.**